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- Definition: prediction of topographic changes in sediment discharges within a computational domain on a specific time window.
- Morphological changes are applied within the following constraints
- 1. Hydrodynamics: conservation of laws of mass and momentum.
- 2. Sediment transport: predictors for sediment transport capacity.
- 3. Bed evolution: conservation law for sediment mass.







- Bedload: particles in a flowing fluid that are transported along the bed, moved by rolling, sliding or saltating.
- Suspended load: portion of the sediment that is carried by a fluid flow which settles slowly enough so that it almost never touches the bed.
- Bed evolution: combination of deposition phenomena of initial conditions, suspended load and bedload.





- Non-cohesive: following equilibrium formulas.
- Cohesive: erosion and deposition laws, consolidation models.
- Mixed-size sediments: moderately/poorly sorted sediment distribution, sand-gravel and sand-mud mixtures.





• Tries to solve the conservative law equation for sediment mass

$$(1-\lambda)\frac{\partial z_b}{\partial t} + \nabla * Q_b = 0$$

 λ : bed porosity

- Q_b : vector of volumetric transport per unit width without pores (m2/s)
- z_b : bed evolution respective to time (m)
- Considers grain sizes between 0.4-29





Bedload transport

• The dimensionless current-induced sediment transport rate

$$\Phi_b = \frac{Q_b}{\sqrt{g(s-1)d^3}}$$

 Bedload formulas are generally computed as function of the Shields number θ:

$$\theta = \frac{\mu \tau_b}{(\rho_s - \rho)gd}$$

- The different bedload transport formulas will depend on the type of sediment that is available.
- Usual classical approach is Meyer-Peter and Mueller formula for Φ_b , considering critical shields parameter equal to 0.047.

 Q_b :bedload transport rate per unit width g: gravity ρ : water density ρ_s : sediment density d: sand grain diameter (= d_{50} for uniform

sediment distribution)

 μ : correction factor for skin friction

 τ_b : bottom shear stress

$$\Phi_{b} = \begin{cases} 0, if \theta < \theta_{cr} \\ \alpha_{mpm}(\theta - \theta_{cr})^{\frac{3}{2}} otherwise \end{cases}$$







• Two-dimensional advection-diffusion equation $\frac{\partial hC}{\partial t} + \frac{\partial hUC}{\partial x} + \frac{\partial hVC}{\partial y} = \frac{\partial}{\partial x} \left(h\varepsilon_s \frac{\partial C}{\partial x}\right) + \frac{\partial}{\partial y} \left(h\varepsilon_s \frac{\partial C}{\partial y}\right) + E - D$

C(x, y, t): depth-averaged concentration expressed in % volume.

- ε_s : turbulent diffusivity of the sediment.
- D: non-cohesive deposition rate
- *E*: non-cohesive erosion rate
- 4 different equilibrium equations are available for solving near-bed concentration are available, depending on the Shields parameter and the diameter of the sediment.



Non-Uniform sediment transport

- In non-uniform bedload sediment transport, moving sediment particles collide and interact.
- Each sediment can be transported by suspended-load or bedload.
- Suspended load mass is exchanged vertically between the water column and the uppermost bed layer.
- Bedload mass is exchanged horizontally between top layer of the bed.

$$\frac{\partial}{\partial t}(hC_k) + \frac{\partial}{\partial x}(hUC_k) + \frac{\partial}{\partial y}(hVC_k)$$
$$= \frac{\partial}{\partial x}\left(h\varepsilon_s\frac{\partial C_k}{\partial x}\right) + \frac{\partial}{\partial y}\left(h\varepsilon_s\frac{\partial C_k}{\partial y}\right)$$
$$+ \omega_{sk}\left(C_{eq_k} - C_{Zrefk}\right)$$





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initial stratification + active layer concept = numerical stratification each layer has its own percentage of sediment class AVAIL(J,K,I) (according to the number of size classes)



Cohesive sediment transport



- Cohesive particles: fine particles like silts and clay that have a diameter less than a limiting value of about 60μ m.
- In cohesive sediment are mainly transported in suspension and transport processes strongly depend on flocculation, consolidation and other physic-chemical processes.
- Similar to before, the transport can be modelled using advection-diffusion equation.

$$\frac{\partial hC}{\partial t} + \frac{\partial hUC}{\partial x} + \frac{\partial hVC}{\partial y} = \frac{\partial}{\partial x} \left(h\varepsilon_s \frac{\partial C}{\partial x} \right) + \frac{\partial}{\partial y} \left(h\varepsilon_s \frac{\partial C}{\partial y} \right) + E - D$$

$$E = \begin{cases} M \left[\left(\frac{\tau_b}{\tau_{ce}} \right) - 1 \right], if \tau_b > \tau_{ce} \\ 0 \\ U = w_s C \left[1 - \left(\frac{\sqrt{\frac{\tau_b}{\rho}}}{u_{mud}^{cr}} \right)^2 \right] \end{cases}$$

$$\frac{dhC}{dt} = W_s C \left[1 - \left(\frac{\sqrt{\frac{\tau_b}{\rho}}}{u_{mud}^{cr}} \right)^2 \right]$$

$$\frac{dhC}{dt} = \frac{dhUC}{dt} \left[\frac{dt}{dt} + \frac{dhUC}{dt} + \frac{dhVC}{dt} + \frac{dhVC$$

Cj the mud concentration of layer j [g/l]



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- Two different sediment classes are considered to solve sediment mixtures
- 1. Non-cohesive sediment: represented by grain diameter D_1 , settling velocity w_{s1}
- 2. Cohesive sediment: represented by grain diameter D_2 that is less than 60 μ m. The setting velocity is a function of the flocs properties which differs from cohesive particles.

$$\frac{\partial hC_k}{\partial t} + \frac{\partial hUC_k}{\partial x} + \frac{\partial hVC_k}{\partial y} = \frac{\partial}{\partial x} \left(h\varepsilon_s \frac{\partial C_k}{\partial x} \right) + \frac{\partial}{\partial y} \left(h\varepsilon_s \frac{\partial C_k}{\partial y} \right) + E^k - D^k$$

$$E_{j} = \begin{cases} M \left[\left(\frac{\tau_{b}}{\tau_{ce}} \right) - 1 \right], & \text{if } \tau_{b} > \tau_{ce} \\ 0 \\ D_{j} = w_{s} C \left[1 - \left(\frac{\sqrt{\frac{\tau_{b}}{\rho}}}{u_{mud}^{cr}} \right)^{2} \right] \end{cases}$$







- Morphological models can be run fully coupled and decoupled.
- Fully coupled: sediment transport and flow occur simultaneously, thus they should be solved at the same time. Examples: debris flow, sediment-laden floods.
- Decoupled: time scale for river or seabed is much longer tan typical time scale for water flow.







- Minimum set of files for running a morphodynamic simulation:
- 1. Steering file(s) (text/ascii file *.cas)
- 2. Geometry file (format selafin/binary *.slf)
- 3. Boundary conditions file (text/ascii file *.cli)
- 4. Additional or optional input files as the fortran file (text/ascii file *.f), reference file (format selafin/binary *.slf)
- * All files should be contained within the same folder.





- Following with a correct outline for morphodynamic simulation, sisyphe's steering file specifies the following:
- 1. Input and output files
- 2. Physical parameters such as sand diameter, settling velocity, etc.
- 3. Main sediment transport processes such as transport mechanisms, closure relationships, etc.
- 4. Additional sediment transport processes like secondary currents, slope effect, etc.
- 5. Numerical options and parameters like numerical schemes, solvers, etc.





/ SISYPHE bedload	
/	/
/ FILES	
/	
/ GEOMETRY	
GEOMETRY FILE	<pre>> '/geo_bifurcation.slf'</pre>
BOUNDARY CONDITIONS FILE /	= '/bc_bifurcation_tel.cli'
/ RESULTS	
RESULTS FILE	= 'res_bifurcation_sis.slf'
/ PHYSICAL PARAMETERS	
/	
/ BED LOAD	= YES
BED-LOAD TRANSPORT FORMULA	= 1
SEDIMENT DIAMETERS	= 0.000120
/	
* * *	
/	/
/ NUMERICAL PARAMETERS	1
/	///////////////////////////////////////
/	
MASS-BALANCE	= YES
SOLVER ACCURACY	= 1.E-12
MASS-LUMPING	= YES
• • •	

- Sediment diameters
- Sediment density
- Sediment parameter (Shields Parameter)
- Settling velocity
- Bed porosity



FORTRAN files

- Just as with telemac2d, FORTRAN subroutines can be specified for the simulation of special conditions using SISYPHE. Examples of these are
- 1. Definition of rigid areas
- 2. New sediment transport formula
- 3. Read data from result files
- Many available subroutines can be found in the folder /sources/sisyphe/ of the TELEMAC-MASCARET SYSTEM.

ime	Date modified	Туре	Size
ad_get_sisyphe.F	1/5/2021 8:45 PM	F File	3 KB
ad_set_sisyphe.F	1/5/2021 8:45 PM	F File	3 KB
ad_sisyphe_checkpoint.f	1/5/2021 8:45 PM	F File	19 KB
Alire	1/5/2021 8:45 PM	File	3 KB
] bedload_bailard.f	1/5/2021 8:45 PM	F File	8 KB
j bedload_bijker.f	1/5/2021 8:45 PM	F File	6 KB
] bedload_calcdw.f	1/5/2021 8:45 PM	F File	5 KB
bedload_dibwat.f	1/5/2021 8:45 PM	F File	8 KB
] bedload_diffin.f	1/5/2021 8:45 PM	F File	9 KB
bedload_direction.f	1/5/2021 8:45 PM	F File	3 KB
] bedload_effpnt.f	1/5/2021 8:45 PM	F File	10 KB
] bedload_einst.f	1/5/2021 8:45 PM	F File	3 KB
bedload_engel.f	1/5/2021 8:45 PM	F File	4 KB
bedload_engel_cc.f	1/5/2021 8:45 PM	F File	4 KB
bedload_evol.f	1/5/2021 8:45 PM	F File	10 KB
bedload_formula.f	1/5/2021 8:45 PM	F File	14 KB
bedload_hiding_factor.f	1/5/2021 8:45 PM	F File	5 KB
bedload_hunz_meyer.f	1/5/2021 8:45 PM	F File	6 KB
bedload_interact.f	1/5/2021 8:45 PM	F File	5 KB





- The boundary condition file would include a number of 13 variables at each node.
- The same variables as for telemac2d apply, with some additional ones
- The flags can be:
- 1. =2: closed boundary (wall)
- 2. =4: free boundary (Nuemann's type)
- 3. =5,6: imposed value(Dirichlet's type).

TELE	MA	C-2	D

LIMBOR LIVEOR MEOR USOR VEOR AUBOR LITEOR TEOR ATEOR ETEOR N K

SISYPHE

THEOR LIGEOR LIVEOR QZEOR USOR VEOR AUBOR LIEBOR/LICEOR EBOR/CEOR ATEOR BIEGR N E

- ✓ LIHBOR: set for water depth
- ✓ LIUBOR: set for discharge or velocity in x direction
- $\checkmark\,$ LIVBOR: set for discharge or velocity in y direction
- ✓ LITBOR: set to tracer
- ✓ LIEBOR: set to bottom elevation
- $\checkmark\,$ LICBOR: set to equilibrium or imposed concentration
- ✓ LIQBOR: set to imposed bedload discharge
- ✓ HBOR: prescribed water depth
- ✓ UBOR: prescribed discharge or velocity in x direction
- ✓ VBOR: prescribed discharge or velocity in y direction
- ✓ AUBOR: friction coefficient on lateral walls
- $\checkmark~$ EBOR: prescribed bed evolution
- ✓ CBOR: prescribed concentration
- ✓ Q2BOR: prescribed bedload discharge





- The variables for graphic output will depend merely on the type of problem that is required to be ran.
- There are specific variables that can be used for each of the problems that can be solved using SISYPHE.

```
U="velocity along a sxis (m/s)";
V="velocity along y axis (m/s)";
C="wawe celerity (m/s)";
H="water depth (m)";
S="free surface elevation (m)";
B="bottom elevation (n)";
F="Froude number";
Q="scalar flowrate of fluid (n2/s)";
I="flowrate along x axis (m2/s)";
J="flowrate along y axis (m2/s)";
M="bed-load discharge (m2/s)";
N="bed-load discharge along x axis (n2/s)";
P="bed-load discharge along y axis (m2/s)";
E="bottom evolution (m)";
R="non erodable bottom";
KS="total bed roughness (m)";
TOB="Bed Shear stress (Totalfriction) (N/m2)";
MU = "Skin friction correction factor";
D50 = "Mean grain diameter";
THETAW="wave angle with axis Oy (deg)";
OSSUSP="suspended load transport rate (m2/s)";
(SBL="bed load transport rate (m2/s)";
W="wave height";
X="wave period";
UWB="wave orbital velocity (m/s)";
1Ai="fraction of sediment of class i in the first layer";
2Ai="fraction of sudiment of class i in the second layer";
kAi="fraction of sediment of class i in the k layer";
kES="thickness of the k layer";
kCONC="concentration of bed layer k";
QSi="bed load transport rate of sediment of class i";
CSi="concentration volumic or mass concentration for class i";
CSAT="saturated concentration (kg/m3)";
A="supplementary variable A";
G="supplementary variable G";
L="supplementary variable L";
O="supplementary variable O"
```



Coupling process

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- SISYPHE can be coupled with both Shallow Water Equations (TELEMAC2D) or Reynolds Averaged Navier Stokes Equations (TELEMAC3D).
- The main function caller is dealt by the steering of either of the latter.
- Hotstart file can be added for a fully developed hydrodynamic model.

INITIAL TIME SET TO ZERO TIME STEP NUMBER OF TIME STEPS	= YES = 20.0 = 100000	1
/ COUPLING WITH SISYPHE		
/ COUPLING WITH SISYPHE STEERING FILE COUPLING PERIOD FOR SISYPHE /	<pre>= 'SISYPHE' = 'run_bifurcation_sis.cas' = 1</pre>	
// INITIAL CONDITIONS		
/ / COMPUTATION CONTINUED PREVIOUS COMPUTATION FILE 	= YES = 'res_bifurcation_hotstart_tel.slf'	8



- If including bed load transport, then the keyword BED LOAD = YES must be included on the steering file.
- Need to select BED-LOAD TRANSPORT FORMULA, which will solve the Shields number depending on the grain size and movement.
- The modification of the bed must consider the following aspects:
- 1. Effect of local bed slope
- 2. Secondary flow effects on direction to bed shear stress
- 3. The bed shear stress partitioning affected by skin friction.



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```
1 : MEYER-PETER and MUELLER
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2 : EINSTEIN-BROWN
```

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3 : ENGELUND-HANSEN + CHOLLET ET CUNGE (total sediment transport)
```

```
30: ENGELUND-HANSEN (total sediment transport)
```

7 : VAN RIJN







Numerical treatments available

- Rigid beds: non erodible beds are treated numerical by limiting bed erosion and letting incoming sediment pass over. This can be modified using the subroutine noerod.f
- Tidal flats: areas where the water depth can become zero during the simulation. Use subroutine positive_depths.f. If selected, can give numerical parameter MINIMUM DEPTH FOR BEDLOAD = 1E-2.
- Morphological factor: increases the bottom change rates with a constant factor N. The keyword MORPHOLOGICAL FACTOR has to be given.
- Sediment slide: preventing bed slope to become greater than the maximum friction angle which is between 32 to 40 degrees. Use subroutine maxslope.f to maintain stability during slides. If option SEDIMENT SLIDE = YES, add friction angle to it.

Useful variables for printout

TOB="Bed shear stress(N/m2)"; MU ="Skin friction coefficient"; M="bed-load discharge (m2/s)"; N="bed-load discharge along x axis (m2/s)"; P="bed-load discharge along y axis (m2/s)"; E="bottom evolution (m)"; QSBL="bed load transport rate (m2/s)";



Suspended sediments transport



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- If including suspended sediment transport, the keyword SUSPENSION = YES must be specified.
- Many concentration formulas are available for near-bed processes, given by keyword REFERENCE CONCENTRATION FORMULA: [1-4]

Diffusion and dispersion for depth-averaged suspended concentration is taken into account with keyword DIFFUSION =YES and OPTION FOR DISPERSION following options

- 1. DISPERSION ALONG THE FLOW, DISPERSION ALONG ACROSS THE FLOW
- 2. Elder model

- 1: Zyserman and Fredsoe
- 2: Bijker
- 3: Van Rijn
- 4: Soulsby & van Rijn

```
TOB="Bed shear stress (N/m2)";
MU="Skin friction coefficient";
M="Solid discharge";
N="Solid discharge along axis x";
P="Solid discharge along axis y";
E="bottom evolution (m)";
QSSUSP="suspended load transport rate (m2/s)";
CSi="concentration volumic or mass
concentration for class i";
```





- If working with sediments that have different sizes, set up the keywords BED LOAD = YES and SUSPENSION = YES.
- Non uniform sediments can be set with keyword NUMBER OF SIZE-CLASSES OF BED MATERIAL from 1 to 20.
- The granulomtric distribution id discretized in the number of classes as a list:
- 1. SEDIMENT DIAMETERS = 0.01
- 2. INITIAL FRACTION FOR PARTICULAR SIZE CLASS = 1, 0, 0...(percentage that must be equal to 1)
- Constant active layers are given with keyword CONSTANT ACTIVE LAYER THICKNESS = number
- For each class, shields critical parameter can be either specified in the cas field or computed by sisyphe using subroutine init_sediment.f

```
E="bottom evolution (m)";
M="bed-load discharge (m2/s)";
lAi="fraction of sediment of class i
in the first layer";
2Ai="fraction of sediment of class i
in the second layer";
kAi="fraction of sediment of class i
in the k layer";
kES="thickness of the k layer";
QSi="bed load transport rate of sediment
of class i"
```

QSSUSP="suspended load transport rate (m2/s)"; CSi="concentration volumic or mass concentration for class i" kCONC="concentration of bed layer k"





- If working with cohesive sediments, with diameters less than 60μ m, specify keyword COHESIVE SEDIMENTS = YES, which sets automatically the following keywords depending on the selected type of sediment: SUSPENSION = YES and BED LOAD = NO
- The number of cohesive layers can be represented as fixed number of layers less than 20 with keyword NUMBER OF LAYERS OF THE CONSOLIDATION MODEL
- Each layer is characterized by its concentration and resistance to erosion. For concentration, use keyword MUD CONCENTRATION PER LAYER =20.; 60.; 100.;... in kg/m3.
- The resistance o each layer can be specified with keyword CRITICAL EROSION SHEAR STRESS OF THE MUD = 0.01.; 0.02.;... in N/m2
- The initialization can also be done with subroutine init_compo_coh.f

kES="thickness of the k layer"; kCONC="concentration of bed layer k"; CSi="concentration volumic or mass concentration for class i";



Mixed sediment transport



• So far it is only supposed suspended load, implying models that solve mixtures of fine sad grains and mud.

Current limitations of TELEMAC2D/3D

- 1. Only one sediment size is allowed for the sand, with constant density for all layers: the volume percentage can vary for different layers.
- 2. Only one sediment size is allowed for the mud: the mass concentration and volume percentage can vary for different layers.
- 3. A simple consolidation of the sand/mud mixture is proposed.

MIXED SEDIMENT = YES

- BED LOAD = NON
- SUSPENSION = YES
- COHESIVE SEDIMENTS = NON

NUMBER OF SIZE-CLASSES OF BED MATERIAL = 2 SEDIMENT DIAMETERS = D1, D2 INITIAL FRACTION FOR PARTICULAR SIZE CLASS = f1, f2 SETTLING VELOCITIES = ws1, ws2 CRITICAL EROSION SHEAR STRESS OF THE MUD= 0.01; 0.02; ...

Initial layer distribution: init_compo_coh.f



Results



